

HYDROLOGICAL OUTLOOK

Brief description of river flow forecasts using seasonal weather forecasts





Funding

The research described in this publication has been funded by the Natural Environment Research Council.

Author

Victoria Bell, Helen Davies, Alison Rudd and Rhian Chapman, Centre for Ecology & Hydrology

Publication Address

Centre for Ecology & Hydrology Maclean Building Benson Lane Crowmarsh Gifford Wallingford Oxfordshire OX10 8BB UK

General and business enquiries: E-mail: +44 (0)1491 692562 enquiries@ceh.ac.uk

Revision date: 11 July 2018

Contents

1	Foreword	4
2	General methodology	.4
3	Technical description: the Grid-to-Grid (G2G) Model	4
4	Technical description: Relative wetness and dryness	. 5
5	Technical description: Monthly forecasts	7
6	Interpretation of results	10
Refe	References	

1 Foreword

This document describes one method of generating a seasonal river flow forecast that is used to inform the UK Hydrological Outlook (Prudhomme *et al.* 2017).

2 General methodology

A methodology for national hydrological outlooks has been developed which combines a hydrological model (Grid-to-Grid or "G2G") estimate of total subsurface water storage (in both soil and groundwater) across Britain with a range of Met Office seasonal rainfall forecasts to provide estimates of area-wide hydrological conditions up to a few months ahead. Bell *et al.* (2013) and Bell *et al.* (2017) provide a summary of the methodology, which is summarised in the technical description below.

For many areas, hydrological forecasts up to a few weeks or months ahead are dependent on accurate knowledge of the current storage of water in the landscape. This information provides the hydrological initial condition, or "initial state", from which future simulations will depart following changes in boundary conditions, consisting primarily of the weather and water consumption. For the Hydrological Outlook UK, the G2G estimate of subsurface water storage provides an initial condition of subsurface water in storage across the UK, derived using the most recent observations of rainfall and potential evaporation (PE).

This current G2G estimate of subsurface water storage not only provides an initial condition for forecasts, it can also provide up-to-date information about how wet or dry the landscape is. Based on this information, maps of "relative wetness" or "relative dryness" can be derived (Section 4) indicating areas where subsurface water storage is particularly high or low, and may be prone to flood or drought conditions in the coming days/weeks. During periods of drought, the link can be made between a deficit in subsurface water storage and a requirement for additional rainfall over subsequent months to enable subsurface water storage and river flow to return to mean monthly values.

The G2G estimated hydrological initial condition provides a starting point from which estimates of water storage and river flows for 1-3 months ahead can be produced as perturbations from the initial state, driven by an ensemble of Met Office rainfall forecasts from the GloSea5 seasonal prediction system. The method used to make these forecasts is summarised in Section 5.

3 Technical description: The Grid-to-Grid (G2G) Model

G2G is a spatially-distributed hydrological model, used in Britain for both continuous simulation of river flows in a changing climate (Bell *et al.*, 2007, 2009) and for real-time flood forecasting (Moore *et al.*, 2006; Cole and Moore, 2009). The model is generally configured to a 1km² grid across the UK, with a 15-minute time-step, and is underpinned by digital spatial datasets on topography, soil/geology and land cover. Gridded time-series of (daily) precipitation and (monthly) potential evaporation are used

as model input and area-wide, gridded time-series of river flows, runoff and soil-moisture are output from the model. Note that the G2G Model simulates natural rather than influenced flows and thus can appear to over- or under-estimate flows for heavily influenced catchments (e.g. those affected by effluent returns or abstractions for public water supply). A detailed description of G2G is presented in Bell *et al.* (2009), with a brief overview of the model's subsurface (soil and groundwater) storage formulation provided below.

For the production of a hydrological outlook, the G2G Model is run continuously over several years to produce an estimate of the most recent hydrological conditions across Britain, from which an estimate is made of the current depth of subsurface water storage as an "anomaly" from climatological mean monthly storage. Here, the depth of subsurface water storage consists of the sum of the unsaturated soil and the groundwater stores.

The soil water state is the volume of available water, V, stored in the unsaturated layer of the soil column of a grid cell of side length Δx . From continuity, $dV/dt = \Delta x^2(P - E - Q)$, with P precipitation, E actual evaporation and Q net outflow per unit area (which includes inflow from upstream cells, lateral flow to the next downstream cell, downward drainage to the saturated zone and saturation-excess surface runoff). Drainage is represented as a simple power law function of V, with two parameters based on soil hydraulic properties derived from the 1km resolution HOST (Hydrology of Soil Types; Boorman *et al.*, 1995) dataset. These data underpin the ability of the model to represent the spatial heterogeneity of subsurface storage: Bell *et al.* (2009) provides more details.

The groundwater state is the volume of available water, V_g , stored in the saturated zone of a grid cell, with drainage from the unsaturated soil column above providing groundwater recharge. A nonlinear function relates groundwater outflow to V_g . Note that while the configuration of soil-storage capacity to the HOST dataset associates an effective maximum to the soil-water volume V in each 1km grid cell, V_g is not limited in this way and its size will reflect the balance between antecedent recharge and

groundwater outflow from the cell. The depth of water in groundwater storage thus arises from the balance between recharge and groundwater outflow over long periods, and while it is unlikely to correspond directly to a groundwater level observation, it can provide an indication of whether storage in the saturated zone is greater or less than the long-term average.

4 Technical description: Relative wetness and dryness

The G2G hydrological model is run continuously over several years to produce an estimate of the most recent hydrological conditions across Britain, from which an estimate is made of the current depth of subsurface water storage (S, mm), consisting of the sum of the unsaturated soil (V) and the groundwater stores (V_g) (mm), i.e. $S = V + V_g$. The example storage anomaly maps in Figures 1 and 2 show daily mean subsurface water storage, S, as an "anomaly" from the historical monthly mean (\bar{S} , mean monthly storage over the period 1981 – 2010). The example storage anomaly map in Figure 1 is predominantly blue indicating a positive anomaly, i.e. with total subsurface water storage greater than the historical monthly mean.

To highlight areas that are particularly wet and dry, the total daily mean subsurface water storage (*S*, mm) is presented using a colour scale showing water storage relative to the historical maximum ("relative wetness", Figure 1) or historical minimum ("relative dryness", Figure 2).



Figure 1 Subsurface storage anomaly and relative wetness for 31st March 2018.



The "relative wetness" map (blue/white map) highlights areas (darker colours) where the G2G simulated subsurface water storage, *S*, approaches or exceeds its historical maximum, S_{max} , for that month. Specifically: the relative wetness is expressed as an anomaly from the monthly mean, $R_w = (S - \bar{S})/(S_{max} - \bar{S})*100$ (%).

Thus a value of R_w >100 indicates the subsurface water storage is higher than the previous maximum monthly value estimated by the G2G over the period 1971 to 2010, and is very wet.

The example storage anomaly map in Figure 2 is predominantly red indicating a negative anomaly. The corresponding "relative dryness" map highlights areas (darker colour) where the G2G simulated subsurface water storage, S, is equal to, or lower than, the historical minimum, S_{min} , for that month.

Specifically: the relative dryness is expressed as an anomaly from the monthly mean,

$$R_d = (\bar{S} - S)/(\bar{S} - S_{min}) *100$$
 (%).

Thus a value of R_d >100 indicates the subsurface water storage is lower than the previous minimum monthly value estimated by the G2G over the period 1971 to 2010, and is very dry.

These maps do not provide a forecast, but the "relative wetness" map can provide an indication of locations which are particularly wet. Rainfall in the high "relative wetness" areas could result in flooding in the coming days/weeks. The "relative dryness" map can provide an indication of locations which are particularly dry, and little or no rain in the high "relative dryness" areas could potentially lead to (or prolong) a drought.

5 Technical description: Monthly forecasts

For the Hydrological Outlook, the G2G provides an initial condition estimate of subsurface water in storage across the UK, derived using the most recent observations of rainfall and PE. This initial condition is a starting point from which estimates of water storage and river flows for 1-3 months ahead can be produced as perturbations from the initial state, driven by Met Office rainfall forecasts.

At present the Met Office can provide an ensemble of weather variables at a seasonal (monthly) resolution (MacLachlan *et al.* 2015, Scaife *et al.* 2014) with a single value for each variable applicable to the whole of the UK. Thus, the high spatial and temporal variability observed in UK weather, particularly rainfall, is not represented. This presents a dilemma for hydrological modellers who typically require high spatial and temporal resolution weather information to estimate a water balance, and to represent the ephemeral nature of streamflow. An ensemble of mean UK rainfall forecasts provides no information on whether the rainfall is more likely to occur in the North or South, however, it does provide some indication of whether the rainfall totals will be higher or lower than the climatological mean (long-term average).

The current generation of Met Office monthly rainfall forecasts can only provide low-resolution temporal and spatial information (UK-scale and monthly). This has led to the development of 1- and 3- month ahead water-balance forecasts of subsurface storage (Bell *et al.* 2013, Bell *et al.* 2017) using UK-wide Met Office forecasts which have been expressed as a perturbation (or "anomaly") from long-term monthly mean rainfall. The use of monthly rainfall anomalies in place of daily or sub-daily rainfall input removes any immediate requirement for temporal downscaling of seasonal rainfall forecasts, which are often provided as monthly totals.

The procedure used to prepare a UK hydrological forecast each month is summarised below:

- 1. The G2G has been run continuously from 1962 to the present day using observed weather data (daily, 1km precipitation and monthly PE data).
- 2. The G2G Model can be initialised by the end-states of a prior run (a scheme commonly used in flood forecasting). Thus the model can be kept up to date by running it forward every month, initialised with the end-states saved at the end of last month, and run forward using as input the latest month's rainfall and PE data.
- 3. The most recent end of month G2G subsurface storage estimate is used as the initial condition for a water-balance forecast of the next 1- and 3-months subsurface storage

using Met Office rainfall forecast ensemble members and climatological PE as input (Bell *et al.* 2013, 2017).

- 4. The Met Office UK-scale monthly and 3-monthly resolution ensemble rainfall forecasts are converted to spatially uniform rainfall anomalies, *a* (mm) relative to Met Office model estimates of climatological mean rainfall. A spatially distributed UK monthly rainfall amount, P^* , is calculated as $P^* = \frac{P_{ij}}{\overline{P}}(\overline{P} + a)$, where \overline{P} and P_{ij} are the UK-mean and the local (1km pixel) monthly mean rainfall (1981-2010) respectively.
- 5. Corresponding ensembles of regional river flow estimates for the next 1- and 3-months ahead can be estimated using a water balance hydrological model (Bell *et al.* 2017).
- Forecast monthly and 3-monthly mean flows are scaled with reference to the long-term monthly mean flow (1981-2010), ranked in relation to 54 years of historical flow estimates (1963 – 2016) and coloured accordingly to one of seven bands, as shown in Figure 1.
- In order to summarise the (~40)-member ensemble of equally-probably hydrological forecasts, 5 maps representing quartiles (4 bands of increasing magnitude) of rainfall (flow) forecasts are presented alongside the maximum and minimum forecasts.
- 8. This type of summary is similar to a box and whiskers summary plot of a distribution, and the resulting set of forecast maps illustrates the median, the inter-quartile range (1st to 3rd quartile) and the overall range (maximum and minimum) of the distribution of forecasts as shown in Figure 1.
- There is a 25% chance of a forecast within each quartile and thus a 50% chance of a forecast within the 1st and 3rd quartiles, as shown in Figure 2.
- 10. Regional histograms (Figure 3) illustrate the full range of ensemble forecasts. They show the percentage of ensemble forecasts falling in each of the flow categories as generated by the monthly-resolution water-balance model. As before results are averaged by region then ranked in terms of 54 years of historical regional flow estimates (1963 2016).



Figure 3. Regional river flow forecasts for June to August 2018 for five members of the Met Office ensemble of rainfall forecasts. The observed flows could be more extreme than the flows generated by either the lowest or highest members of the rainfall ensemble.



Figure 4. The probability of a forecast lying within the quartile bands.



Figure 5. Regional histograms illustrating the full range of ensemble forecasts (June and June-July-August 2018).

6 Interpretation of results

For the examples shown in Figures 3 to 5, the forecasts can be interpreted as follows:

SUMMARY: During June river flows are likely to be in the *Normal range* or below, except in Severn Trent and Anglian where flows are likely to be in the *Normal range* or above.

Over the next 3 months river flows are likely to be in the *Normal range* or below.

Typically, much of the spatial variability (~60%) in the 1-month ahead flow estimates arises from the hydrological initial condition at the start of the month as estimated by the G2G Model, combined with a model estimate of how much water the landscape can store (Bell *et al.*, 2017). For 3-month ahead forecasts, more of the forecast skill arises from the GloSea5 rainfall forecasts (~70%), particularly in the North and West of Britain. Note that any spatial variability arising from the use of the GloSea5 rainfall forecasts comes only from the imposed monthly mean UK rainfall distribution (Section 5, procedure item 4).

The hindcast analysis by Bell *et al.* (2017) also indicated that only very limited forecast skill is achievable for spring and summer seasonal hydrological forecasts driven by GloSea5 rainfall forecasts; however, autumn and winter flows can be reasonably well forecast using (ensemble mean) rainfall forecasts from GloSea5, particularly for 3-months ahead.

7 References

Bell VA, Davies HN, Kay AL, Brookshaw A, Scaife, AA. 2017. A national-scale seasonal hydrological forecast system: development and evaluation over Britain, *Hydrol. Earth Syst. Sci.*, **21**, 4681-4691, doi:10.5194/hess-21-4681-2017.

Bell VA, Davies HN, Kay AL, Marsh TJ, Brookshaw A, Jenkins A. 2013. Developing a large-scale waterbalance approach to seasonal forecasting: application to the 2012 drought in Britain. *Hydrological Processes*, **27**(20), 3003-3012. doi:10.1002/hyp.9863.

Bell VA, Kay, AL, Jones, RG, Moore, RJ. 2007. Use of a grid-based hydrological model and regional climate model outputs to assess changing flood risk. *Int. J. Climatol.*, **27**: 1657-1671. DOI:10.1002/joc.1539.

Bell VA, Kay AL, Jones RG, Moore RJ, Reynard NS. 2009. Use of soil data in a grid-based hydrological model to estimate spatial variation in changing flood risk across the UK. *Journal of Hydrology* **377**(3–4): 335–350.

Boorman DB, Hollis, JM, Lilly A. 1995. Hydrology of soil types: a hydrologically based classification of the soils of the United Kingdom. *IH Report No. 126*, Institute of Hydrology, Wallingford, UK, 137pp.

Cole SJ, Moore, RJ. 2009. Distributed hydrological modelling using weather radar in gauged and ungauged basins. *Advances in Water Resources* **32**(7): 1107-1120. DOI: 10.1016/j.advwatres.2009.01.006

MacLachlan C, Arribas A, Peterson KA, Maidens A, Fereday D, Scaife AA, Gordon M, Vellinga M, Williams A, Comer RE, Camp J and Xavier P, 2015. Description of GloSea5: the Met Office high resolution seasonal forecast system. *Q. J. R. Met. Soc.*, DOI: 10.1002/qj.2396.

Moore RJ, Cole SJ, Bell VA, Jones DA. 2006. Issues in flood forecasting: ungauged basins, extreme floods and uncertainty. In: I. Tchiguirinskaia, K. N. N. Thein & P. Hubert (eds.), *Frontiers in Flood Research*, 8th Kovacs Colloquium, UNESCO, Paris, June/July 2006, IAHS Publ. 305, 103-122.

Prudhomme C, Hannaford J, Boorman D, Knight J, Bell V, Jackson C, Svensson C, Parry S, Bachiller-Jareno N, Davies HN, Davis R, Harrigan S, Mackay J, McKenzie A, Rudd AC, Smith K, Ward R, Jenkins A. 2017. Hydrological Outlook UK: an operational streamflow and groundwater level forecasting system at monthly to seasonal time scales. *Hydrological Sciences Journal* **62**(16), 2753–2768. https://doi.org/10.1080/02626667.2017.1395032.

Scaife A.A., Arribas A, Blockley E, Brookshaw A, Clark RT, Dunstone N, Eade R, Fereday D, Folland CK, Gordon M, Hermanson L, Knight JR, Lea DJ, MacLachlan C, Maidens A, Martin M, Peterson AK, Smith D, Vellinga M, Wallace E, Waters J and Williams A, 2014. Skilful Long Range Prediction of European and North American Winters. *Geophys. Res. Lett.*, **41**, 2514-2519. DOI:10.1002/2014GL059637.